

# Geophysical Tests for Habitability in Icy Ocean Worlds

Steve Vance 5/23/17

Confidential manuscript submitted to JGR Planets

#### Geophysical tests for habitability in ice-covered ocean worlds

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Vance et al., in submitted.

#### Icy and Exo Ocean Worlds

Advances in computational capabilities enable **new equations of state** based on experiments

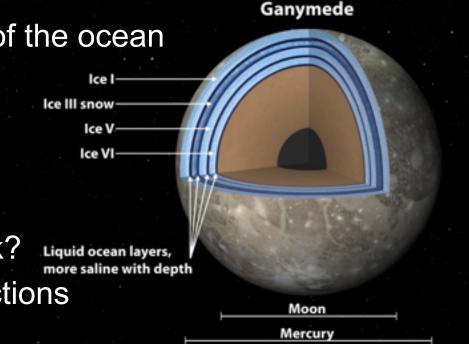
Requiring self-consistent thermodynamics and accounting for ocean salinity affects **how extraterrestrial oceans work**:



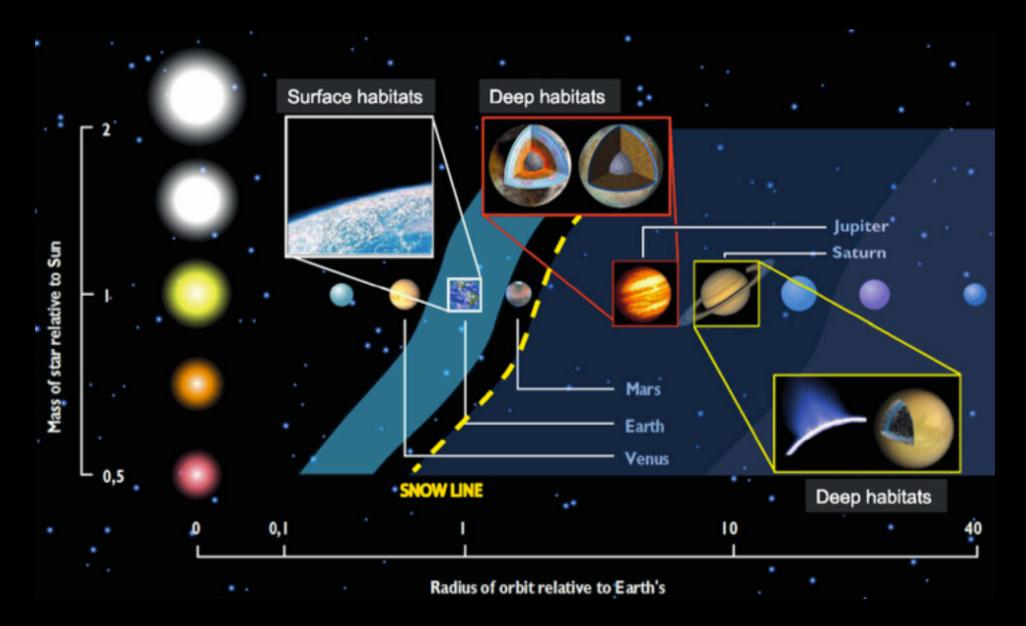
-reduces presence of ices

-lead to layered oceans-ices

—> when does ice float or sink? —> how might water-rock reactions create food for life



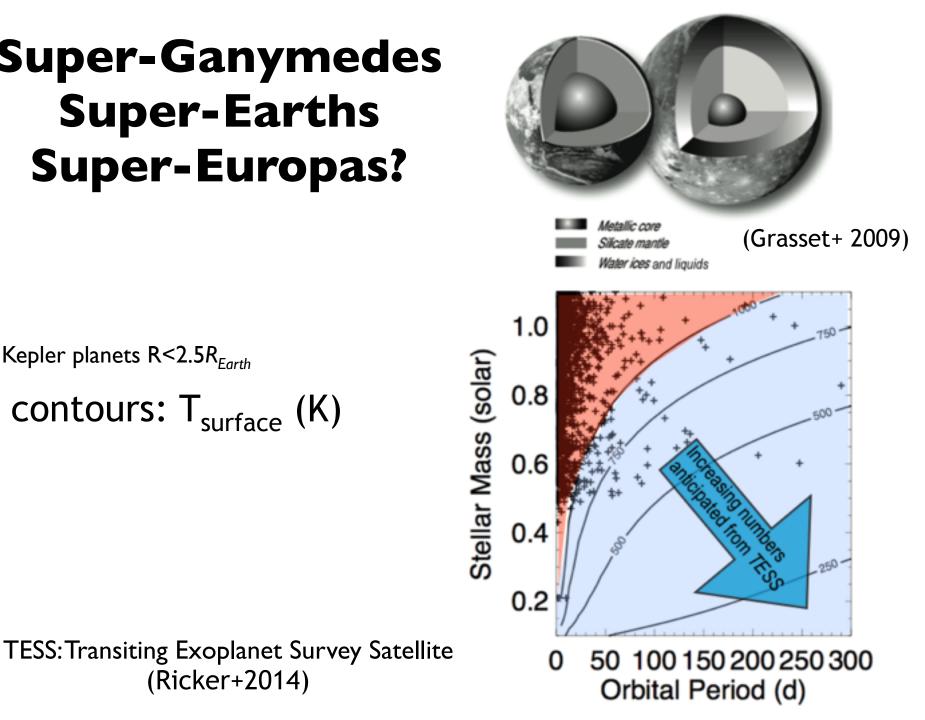
#### Icy worlds expand the "habitable zone".



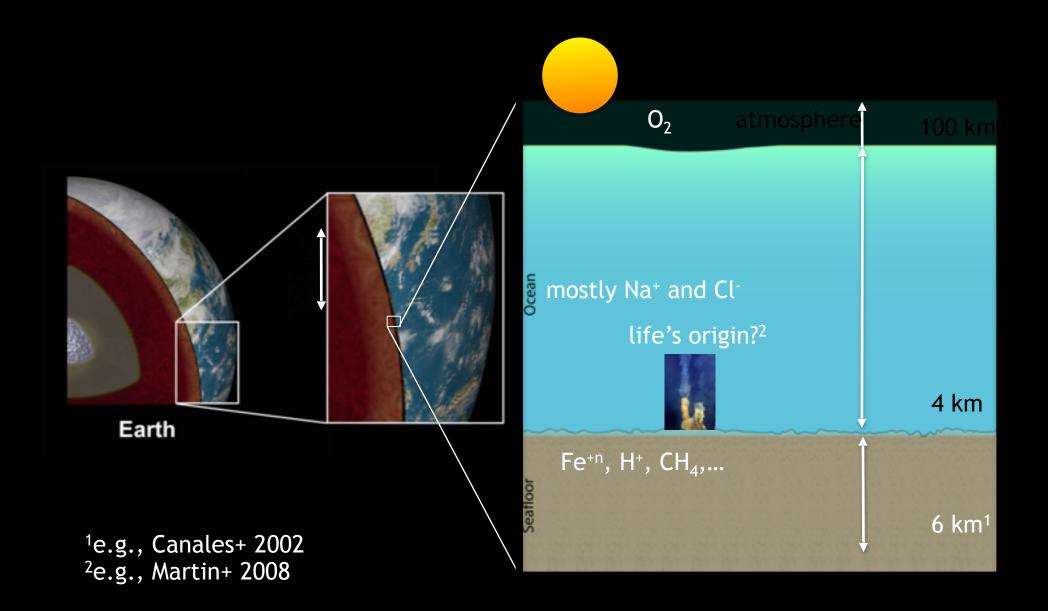
#### **Super-Ganymedes Super-Earths Super-Europas?**

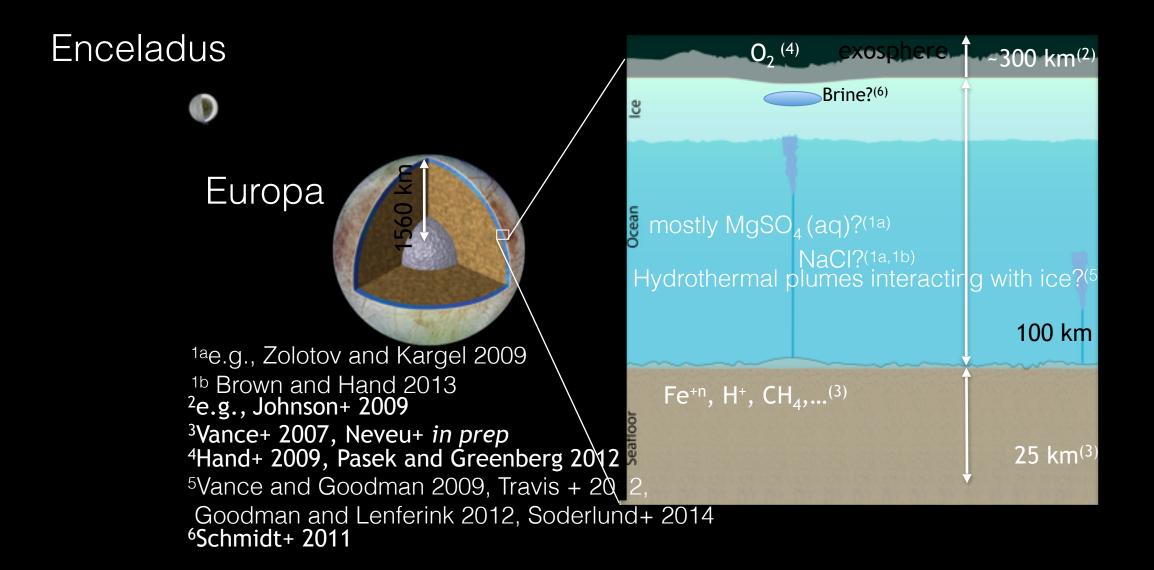
Kepler planets R<2.5R<sub>Earth</sub> contours: T<sub>surface</sub> (K)

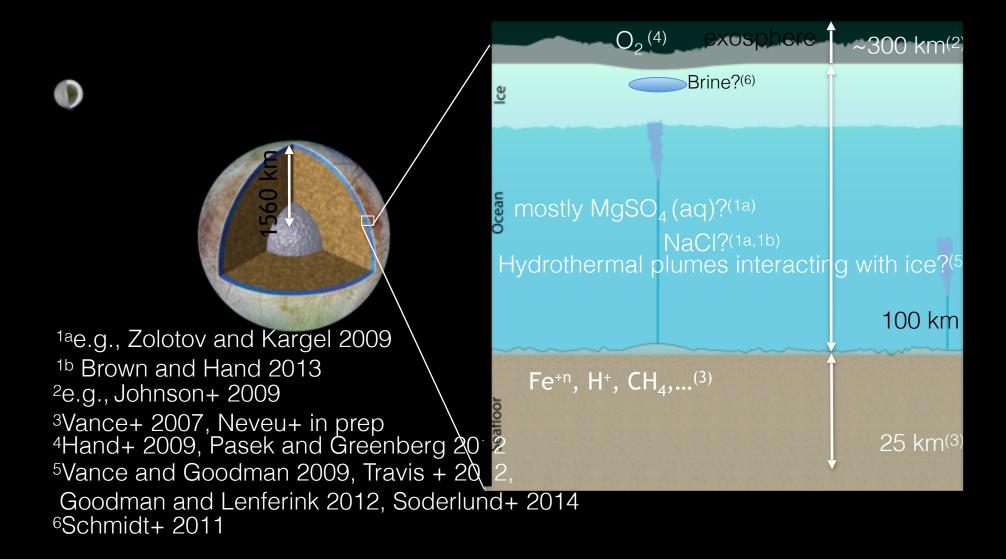
(Ricker+2014)

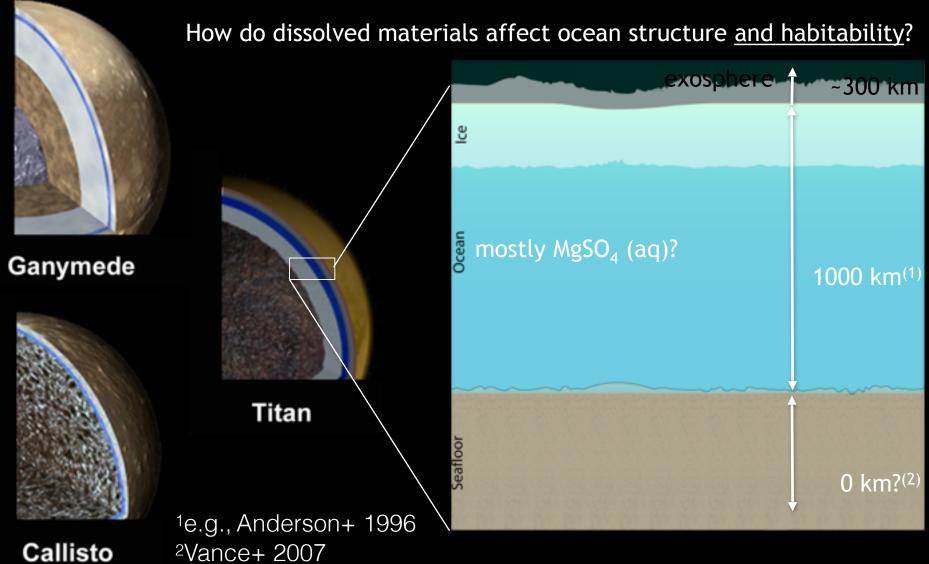


#### How do extraterrestrial oceans compare with Earth's?



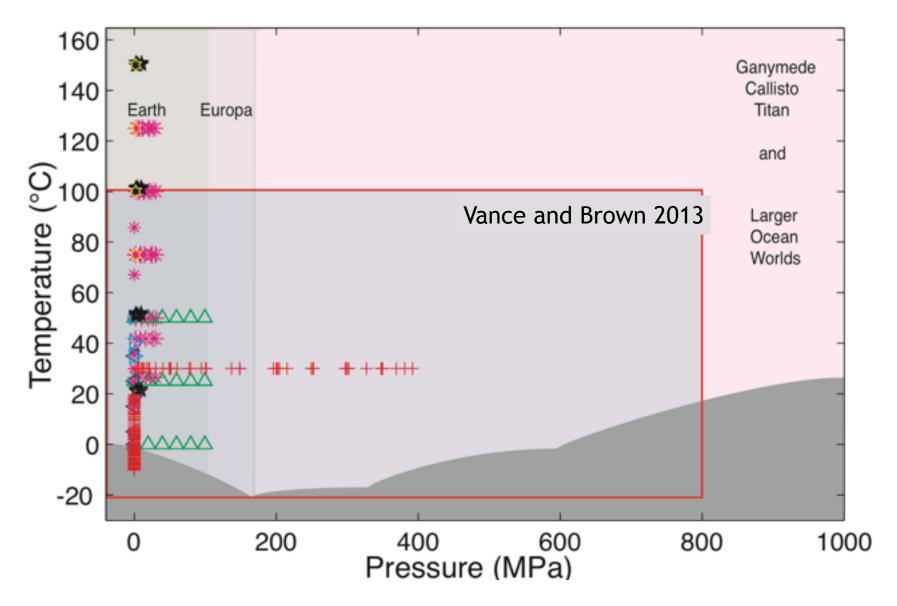


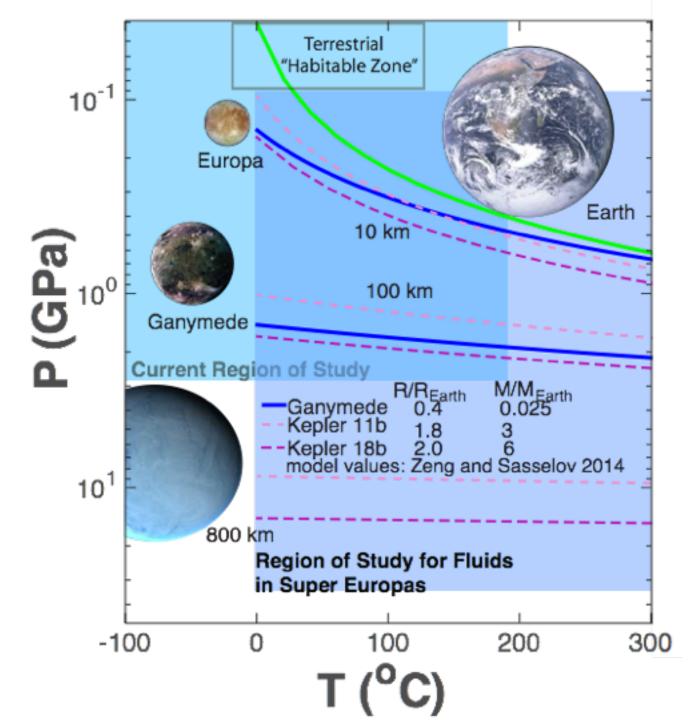




Callisto

#### Knowledge of MgSO<sub>4</sub>(aq) fluid properties

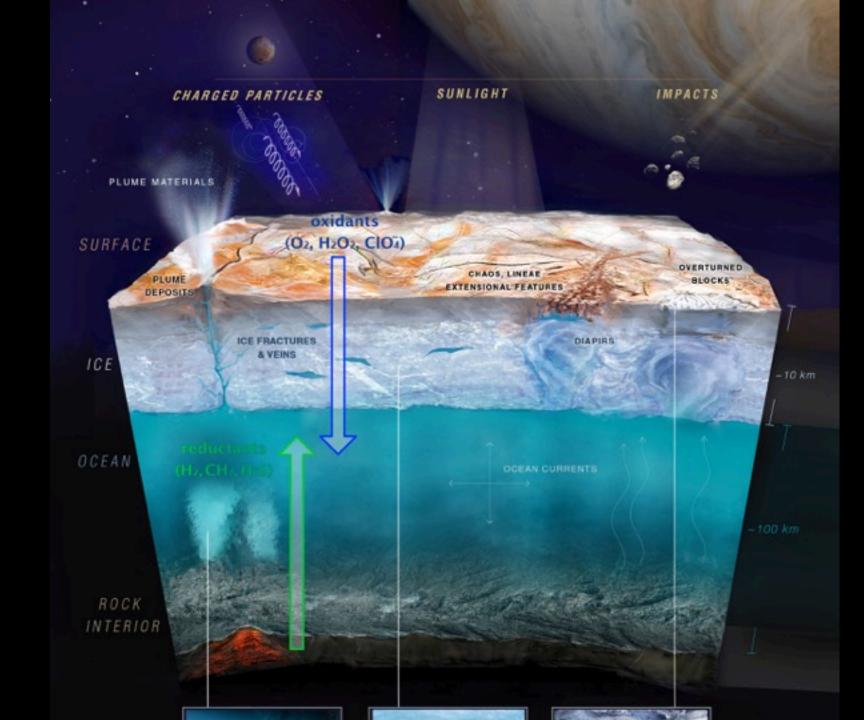


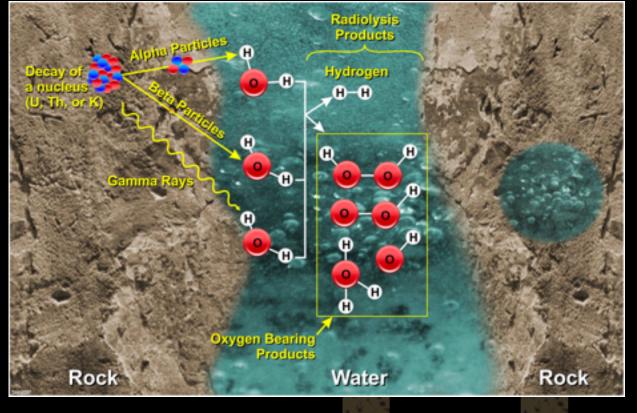


# How can we use geophysical measurements to test habitability?

## Europa

Modified from the 20<u>16</u> Lander SDT Report





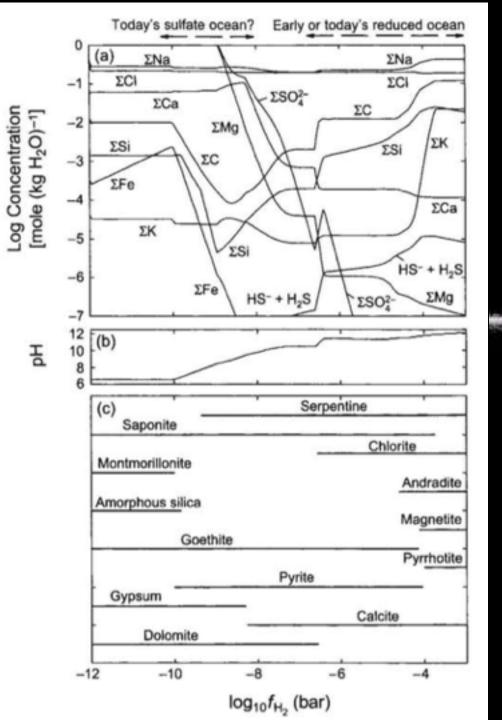
Bouquet et al. 2017

0,-Enceladus H2,CH4 Fe2+→Fe3+ Vance et al. 2007, 2016

Europa



All Purpose Clean

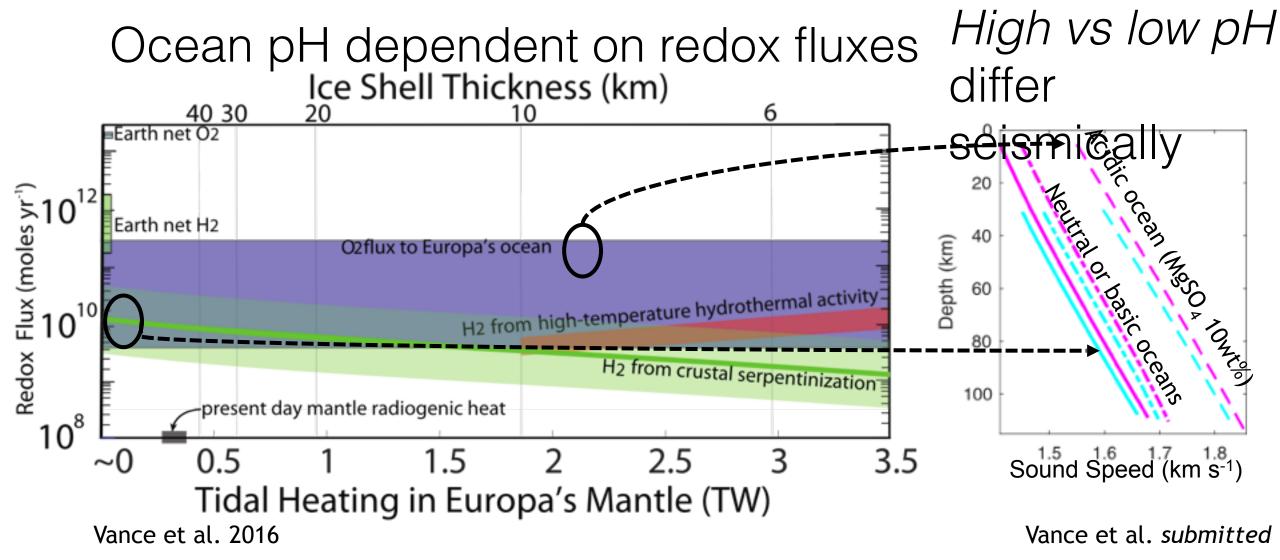




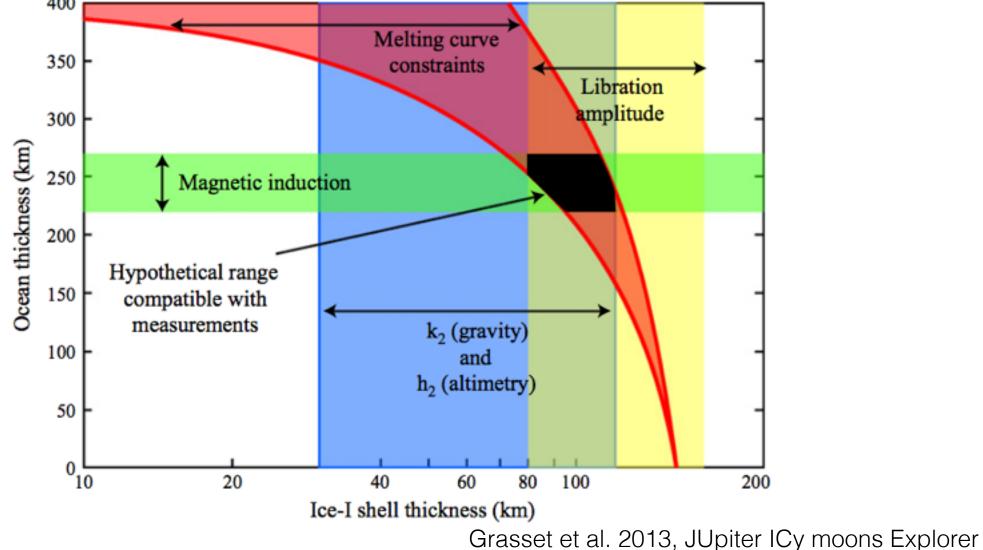
NaCl

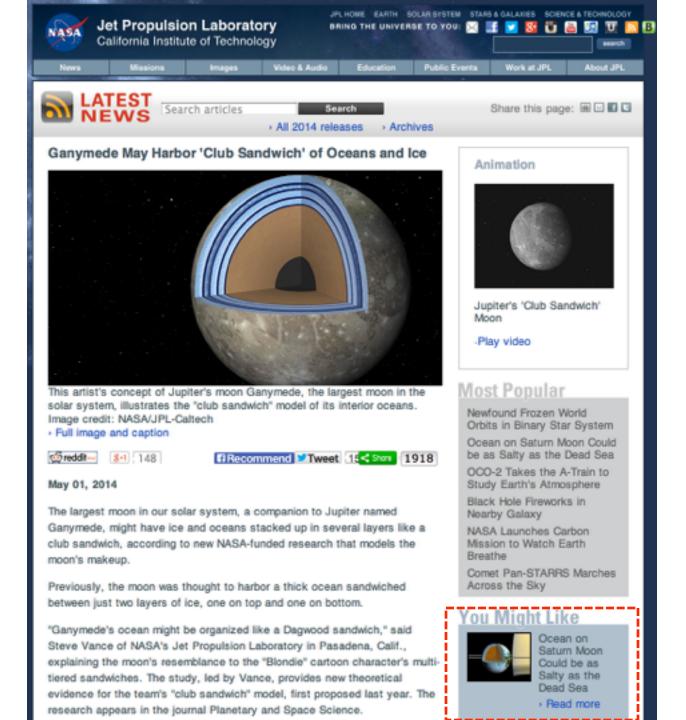
#### Zolotov and Kargel 2009

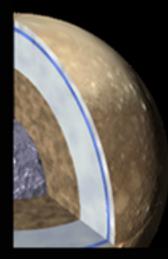
# Seismological Constraints on Habitability



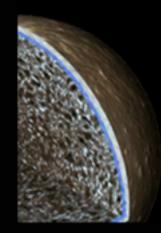
### Geophysical measurements (e.g. at Ganymed







Ganymede



# Titan

boundary condition  $T_b$  of ice I

Seafloo

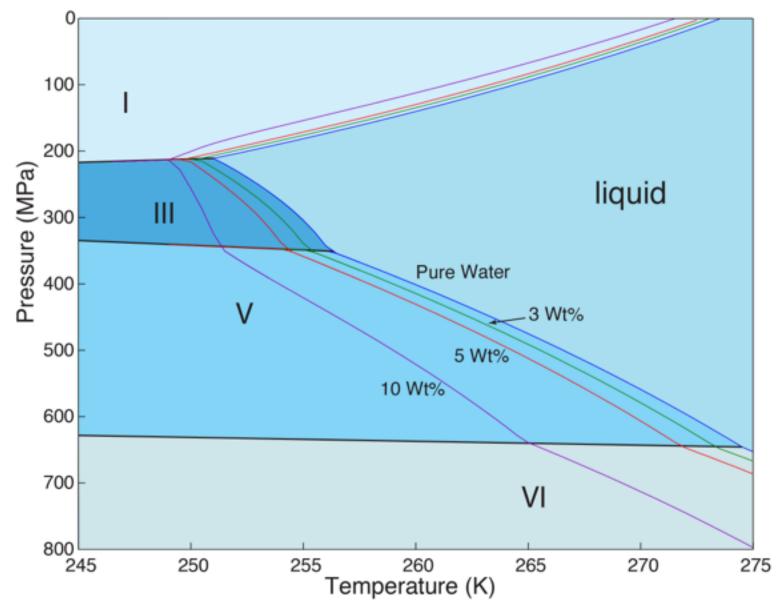
 $\underline{\mathfrak{B}}$  conductive, k=D/T  $\rightarrow T(z) = T_b^{z/z_b} T_o^{1-z/z_b}$ 

 $q_b = D \frac{\ln(T_b/T_s)}{z_b}$ convective, uniform salinity  $\frac{\partial P}{\partial T} = \frac{\alpha T}{\rho C_P}$ 

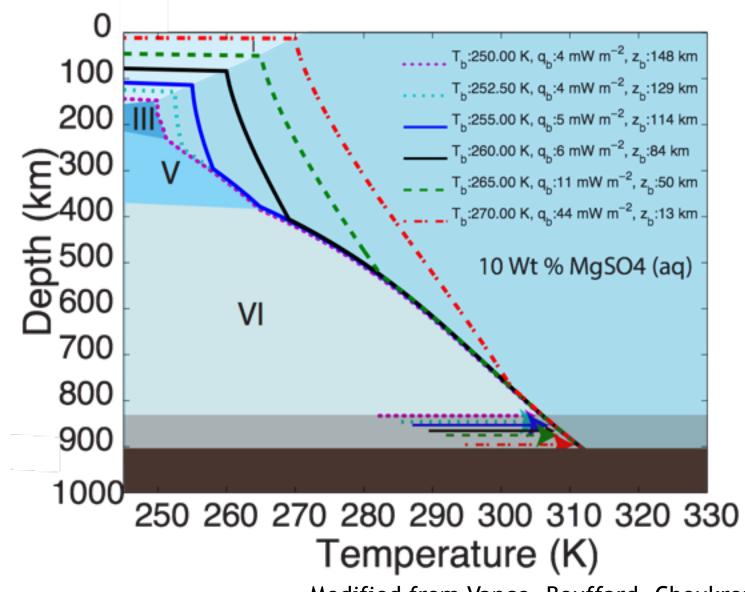
boundaries determine thicknesses

Callisto

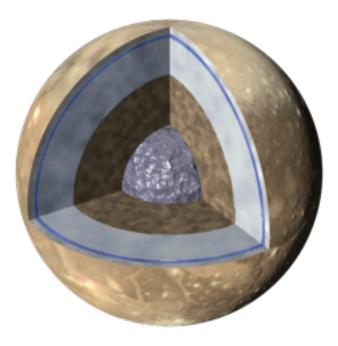
Phase diagram for aqueous magnesium sulfate (MgSO<sub>4</sub>)



Vance, Bouffard, Choukroun, and Sotin, 2014, PSS



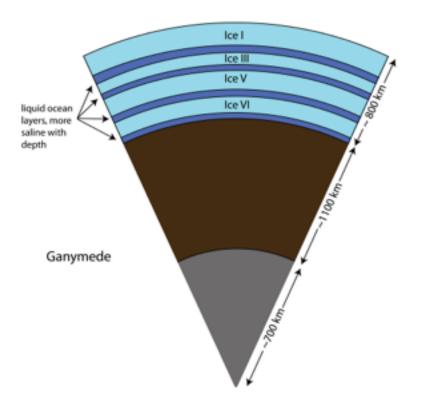
.,



Modified from Vance, Bouffard, Choukroun, and Sotin, 2014,

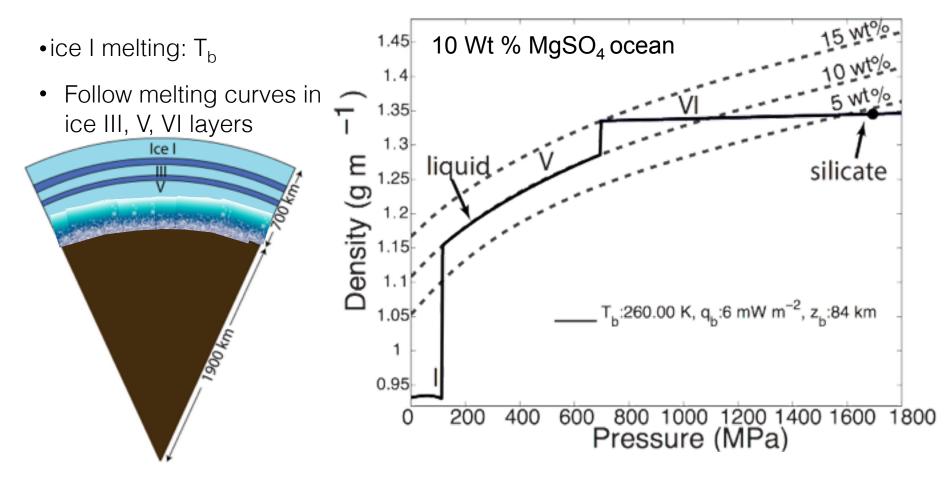
#### Thermal model constrains Ganymede's internal structure

Calculate radii from Galileo constraints on Ganymede's density (1,942±4.8 kg/m3) and gravitational moment of inertia (C/MR<sup>2</sup> = 0.3105±0.028) Schubert+ 2004



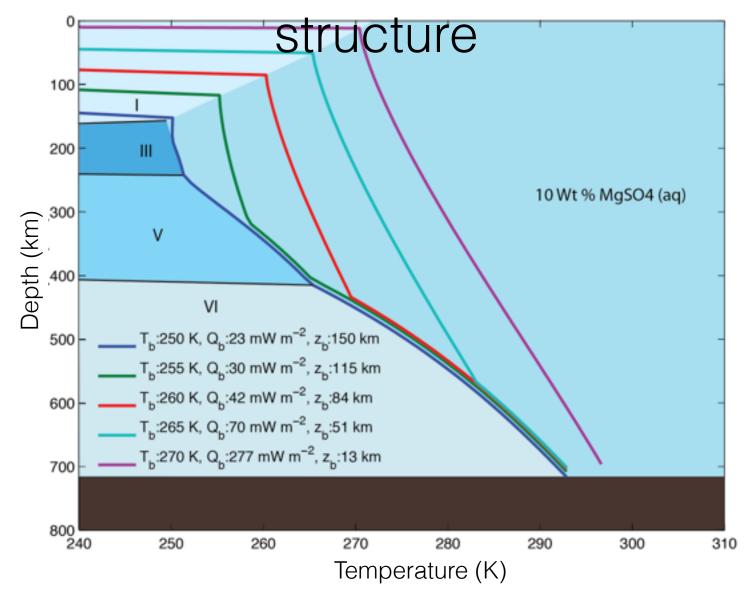
Vance, Bouffard, Choukroun, and Sotin, 2014, PSS

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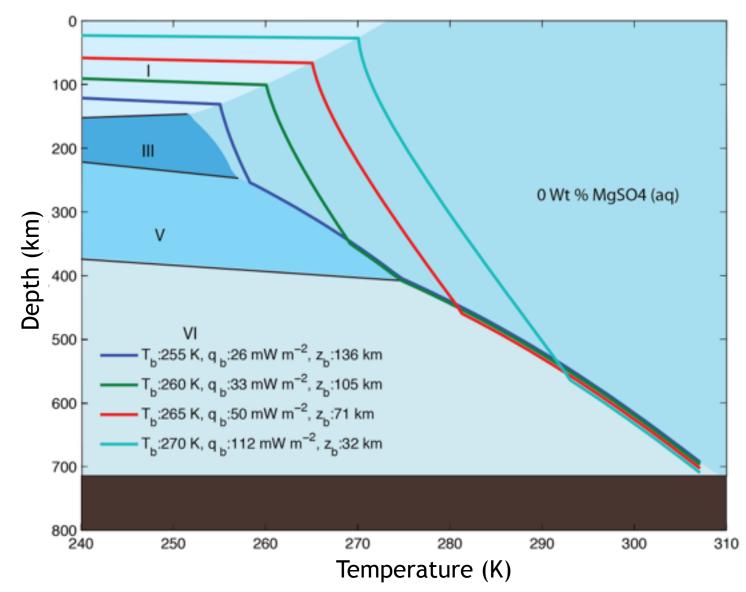
Modified from Vance, Bouffard, Choukroun, and Sotin, 2014,

Salinity determines cryosphere



Vance, Bouffard, Choukroun, and Sotin, PSS 2014.

#### Salinity determines cryosphere



Vance, Bouffard, Choukroun, and Sotin, PSS 2014.

Europa Enceladus Genymede Callisto Titan								if including an Fe core	
ıts	R	ρ	C/MF <sup>e</sup>	To	T <sub>b</sub>	w	Q	ρ <sub>sii</sub>	ρ <sub>Fe</sub>
Input	Radius	Bulk Density	Gravitational Moment of Inertia	Surface Temp.	Ice I bottom Temp.	Ocean Salinity	Mantle Heat	Silicate Mantle Density	Iron Core Density

#### PlanetProfile

github.com/vancesteven/PlanetProfile Vance et al. *submitted.* 

	Radius (km)	Density (kg $m^{-3}$ )	Moment of Inertia
Europa	$1565.0{\pm}8.0$	$2989 {\pm} 46$	$0.346 {\pm} 0.005$
Ganymede	$2631{\pm}1.7$	$1942.0{\pm}4.8$	$0.3115{\pm}0.0028$
Callisto	$2410.3 \pm 1.5$	$1834.4 \pm 3.4$	$0.3549{\pm}0.0042$
Enceladus	$252.1{\pm}0.2$	$1609 \pm 5$	0.335
Titan	$2574.73 {\pm} 0.09$	$1879.8 {\pm} 0.004$	$0.3438{\pm}0.0005$

Schubert et al. 2004; less et al. 2010, 2012

# Fluid

#### phase equilibria, α,ρ,C<sub>P</sub>,ν, G<sub>S</sub>,K<sub>S</sub>

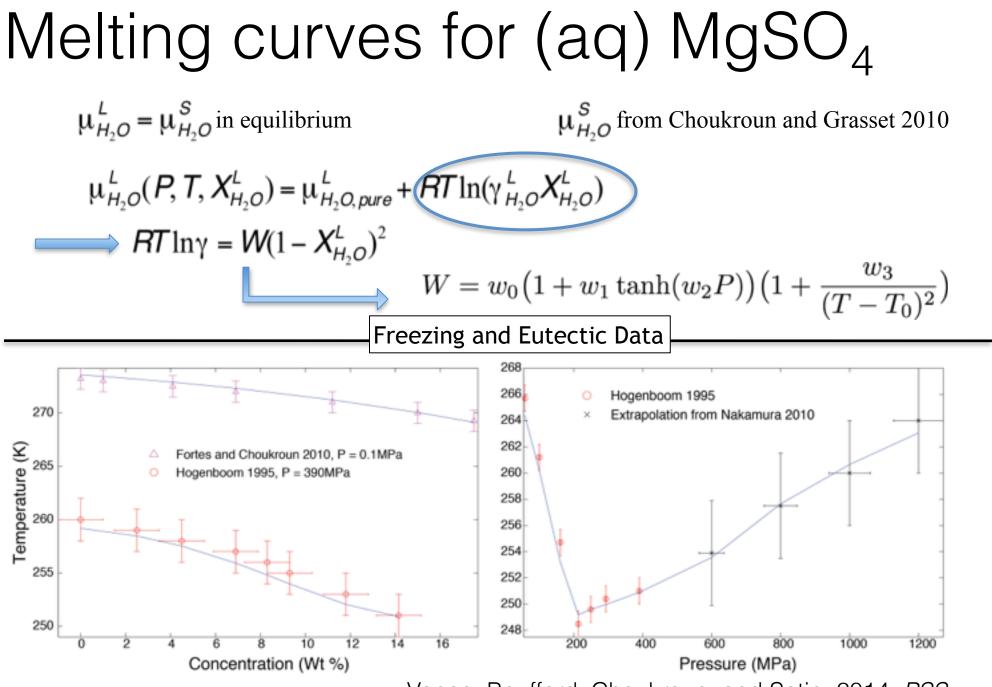
- MgSO<sub>4</sub> (aq) (Vance and Brown 2013, Vance et al. 2014)
  - 0-2 molal; 250-400 K; 0-2 GPa
- Seawater (NaCI) Gibbs Seawater (McDougall 2011)
  - 0-2S0; 250-400 K; 0-0.1 GPa
- NH<sub>3</sub> (aq) (Tillner-Roth and Friend 1998, Choukroun et al. 2010)
  - 0-10 wt%; 250-400 K; 0-2 Gpa

#### lce

Ice Ih, II, III, V, VI (Choukroun et al. 2010)

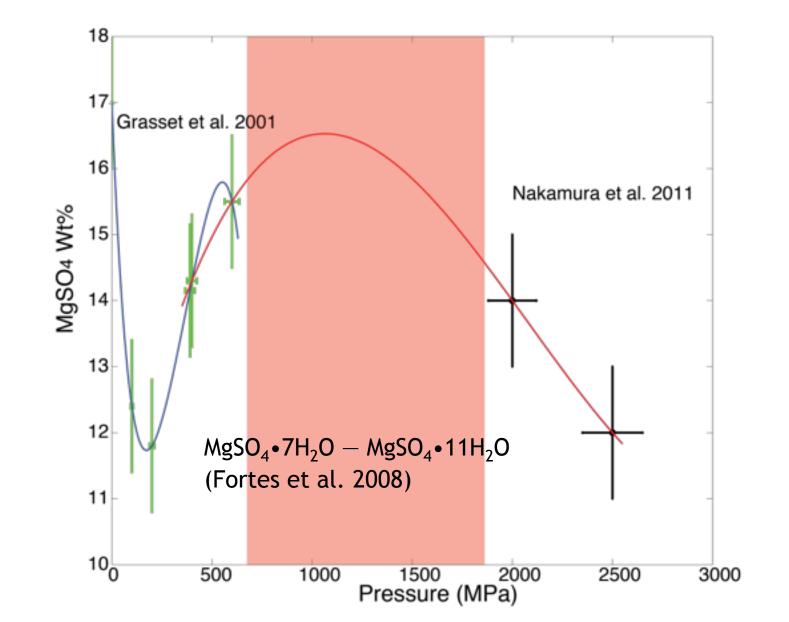
Margules activity model for melting can be replaced with chemical potentials for ice and water being developed by the NAI Icy Worlds team (Brown et al....)

#### Rock Perple\_X (Connolly et al. 2009) using databases from Holland and Powell (2011), Stixrude and Lithgow-Bertollini (2011)



Vance, Bouffard, Choukroun, and Sotin, 2014, PSS

#### Eutectic composition vs pressure



# Melting curves for (aq) NH<sub>3</sub>

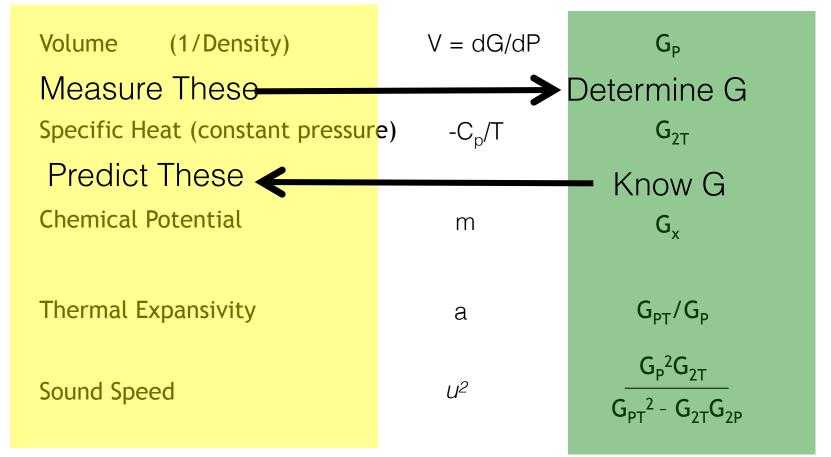
• Activity model for solid-liquid stability:

$$= W(1 - X_{H_2O}^L)^2 W = w_0 (1 + w_1 \tanh(w_2P)) (1 + w_3T + \frac{w_4}{T})$$

Choukroun and Grasset 2010

#### Gibbs Energy Derivatives

G is solution of an ODE

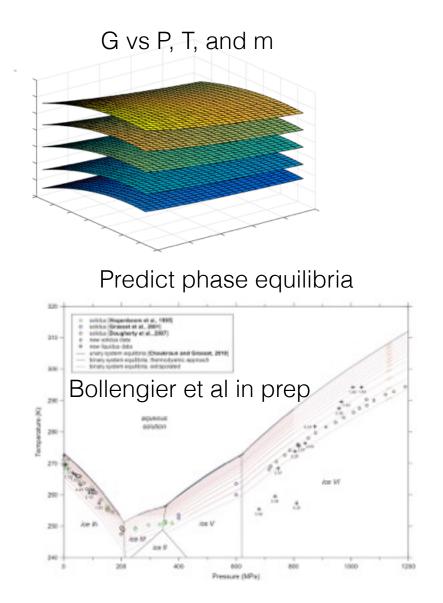


Gibbs energy at high pressure is accurately determined from sound speeds vs P, T, and X

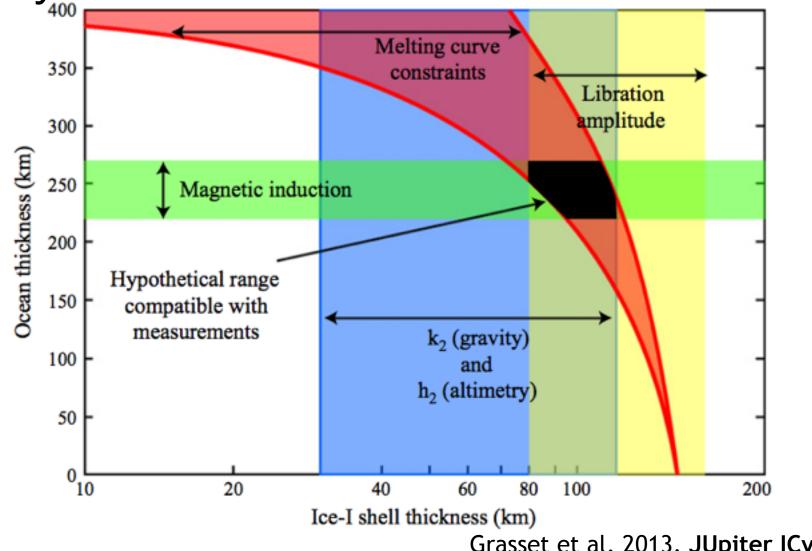
#### Workflow for EOS measurements dict thermodynamic properties

Collect lab measurements in fluids and fit sound speeds

2 1.5 ofy (km s<sup>-1</sup>) 200 300 400 500 600 700 800 Temperature (\* C) Pressure (MPa) Prossure(MPa) Vance and Brown 2013



#### Geophysical measurements

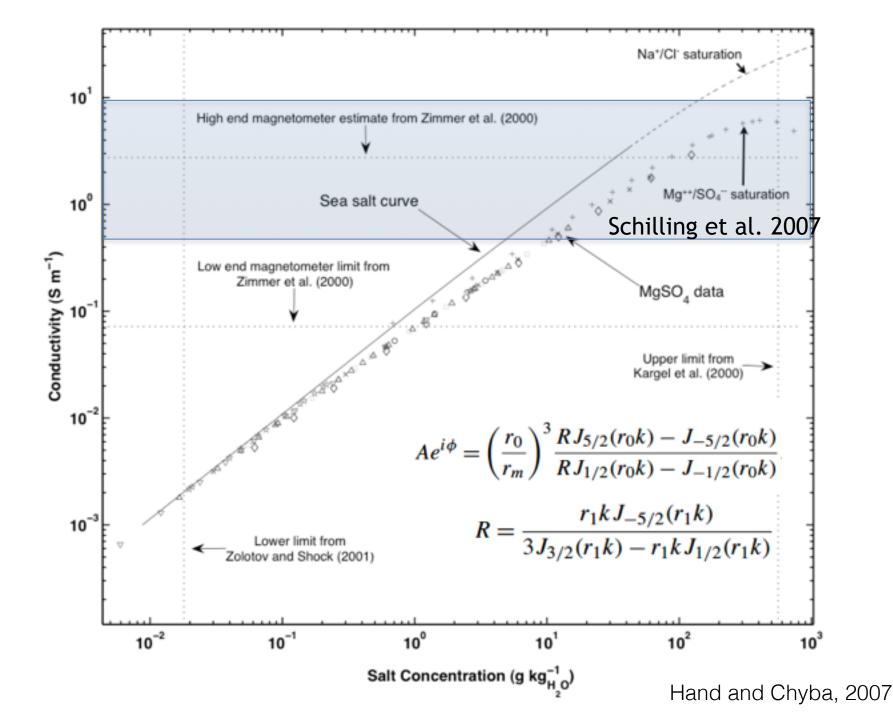


Grasset et al. 2013, JUpiter ICy moons Explorer

# Salinity

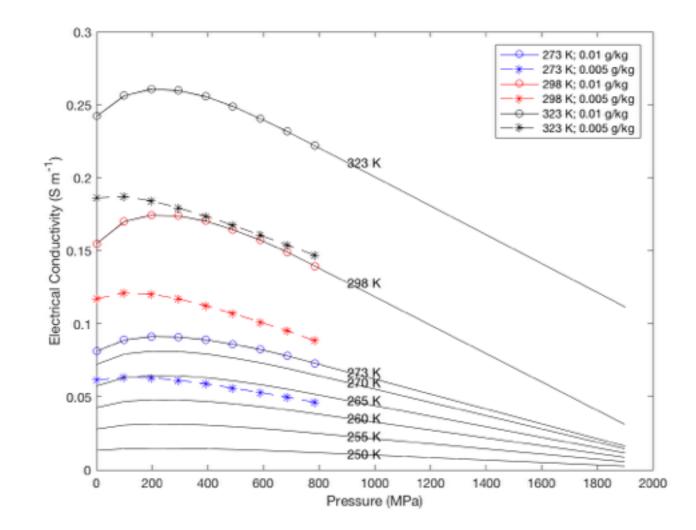
## Salinity High pH - NaCl

Seawater salinity, mostly NaCl, Available from TEOS-10 0 to 42 psu (~g/kg) -35 to 2 °C 0 to 100 MPa Hill 1986





- Available for MgSO<sub>4</sub>: 0.005 and 0.01 moles / kgH<sub>2</sub>O 0.1 to 800 MPa 298 to 423 K
- Larionov and Kryukov (1984)
- Extrapolated to high concentration and low temperature



	Pyrolite <sup>a</sup>	L-LL Chondrite <sup><math>b</math></sup>	Pyrolite <sup>b</sup>	L-LL Chondrite <sup>b</sup>	
SiO <sub>2</sub>	38.71	38.73	38.66	42.78	
MgO	49.85	38.73	48.53	39.68	
ons	6.17	14.98	5.72	13.98	
CaO	2.94	2.12	3.50	2.13	
Al <sub>2</sub> O <sub>3</sub>	2.22	1.36	3.59	1.43	
Na <sub>2</sub> O	0.11	1.02	-	-	

<sup>a</sup>Cammarano et al. [2011];<sup>b</sup>modified from Cammarano et al. [2006]

#### Core composition and properties

	fcc-y-Iron <sup>a</sup>	Fe-S (5% < $X_{FeS}$ < 20%) <sup>b</sup>		
ρ	8000	5150-(%S-10)×50		
$\alpha$ , 10 <sup>-5</sup> K <sup>-1</sup>	5	9.2		
K <sub>S</sub> , GPa	156	53.2-(%S-10)×2		
$\partial K_S / \partial P$	5.0	4.66		
$\partial K_S / \partial T$ , Pa K <sup>-1</sup>	-0.040	-		
G, GPA	76.5	-		
$\partial G/\partial P$	2	-		
$\partial G/\partial T$ , Pa K <sup>-1</sup>	-0.023	-		

<sup>a</sup>Cammarano et al. [2006];<sup>b</sup>Sanloup et al. [2000]

Submitted work Considers historical "pyrolite" and "chondrite" composition as per Kuskov and Kronrod (2001) Adds Na<sub>2</sub>O 1wt% and saturated hydration

Core is either gamma iron or variable FeS

Future work can explore variable mineralogy. e.g.:

mantle iron, MgO/(MgO+FeO) sulfur partitioning into core and ocean

## Elastic Properties

## Elastic Properties - Ices

Crystalline	$K_S$	$K'_S$	$K_S''$	$\mu$	$\mu'$	$\mu^{\prime\prime}$
Phase	GPa		$GPa^{-1}$	GPa		$GPa^{-1}$
$I_h$	9.5	0.33	-0.026	3.3	0.537	025
II	13.89	1.6		5.15	3.5	
III	8.9	3.65		2.7	6.55	
V	11.8	4.8		5.7	0.9	
VI	14.6	4.1		5.0	3.0	

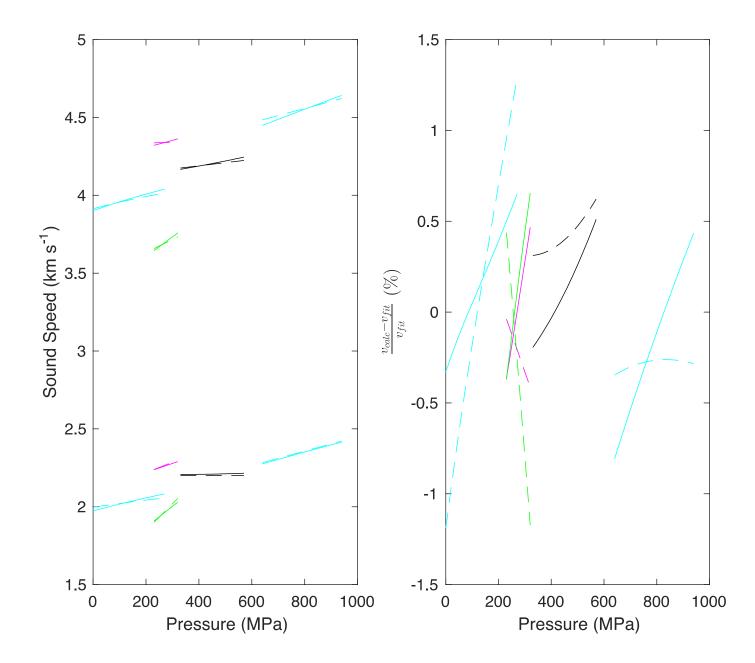
 $V_S = \left[\frac{\mu}{\rho}\right]^{1/2}$  $V_P = \left[\frac{K_S}{\rho} + \frac{4}{3}V_S^2\right]^{1/2}$ 

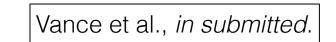
Sound speed in ice based on Shaw 1986 and Gagnon et al. 1988, 1990 from fits to measurements around -35°C

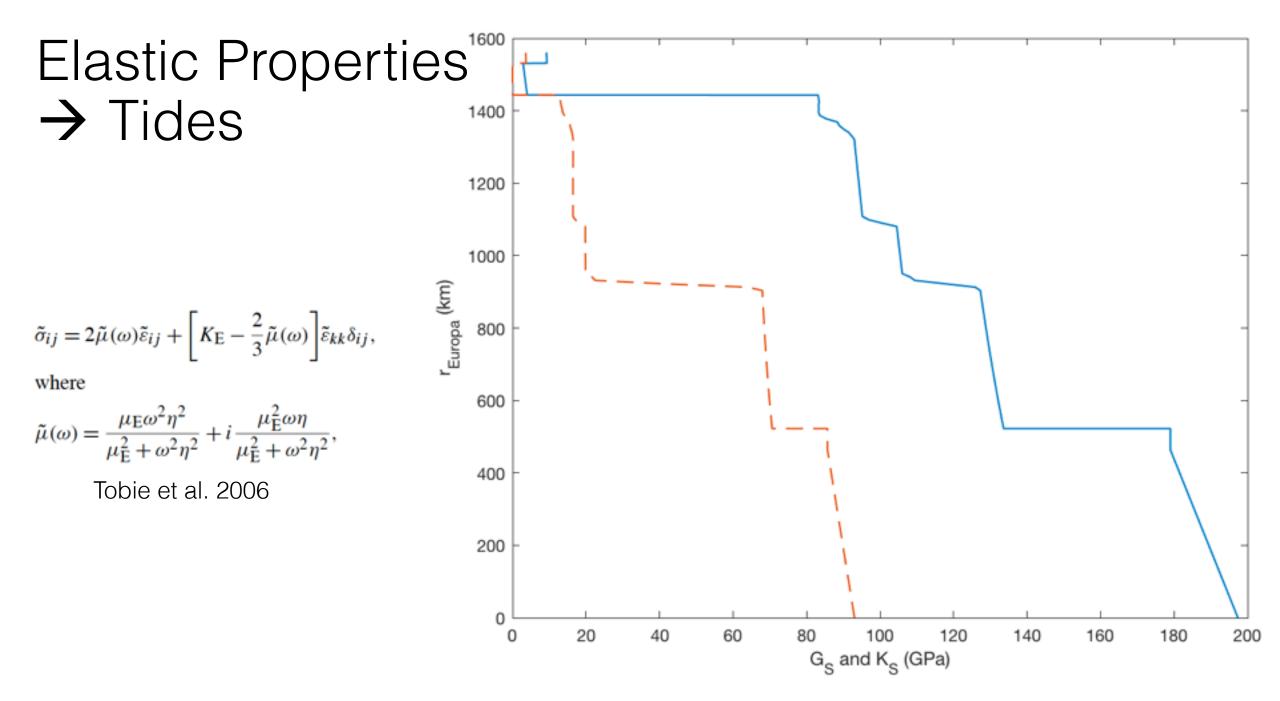
forward model for bulk sound speed from ice equations of state based on density and phase boundaries

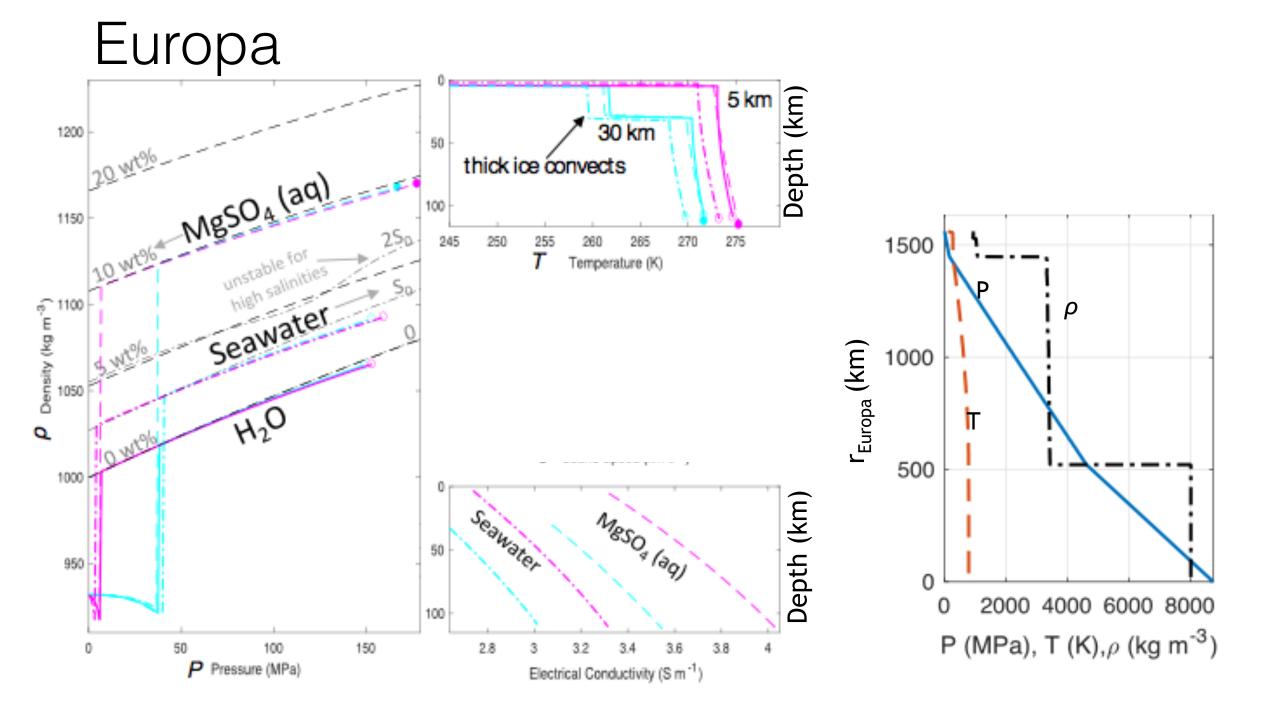
$$\frac{\partial \rho}{\partial P_T} = \frac{1}{v_c^2}$$

Choukroun et al. 2010

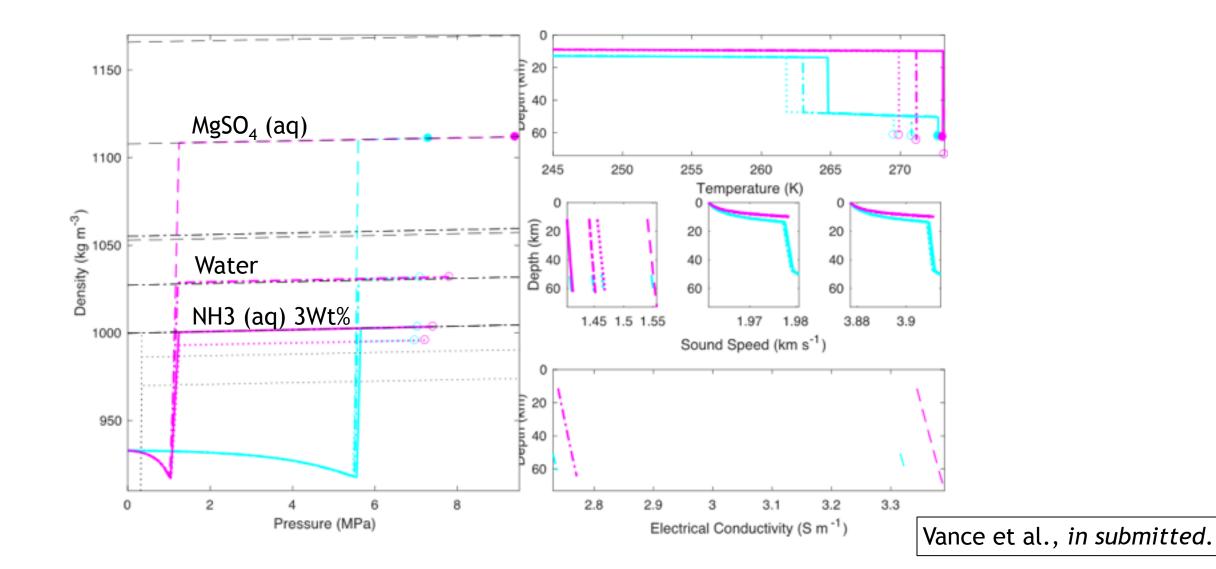






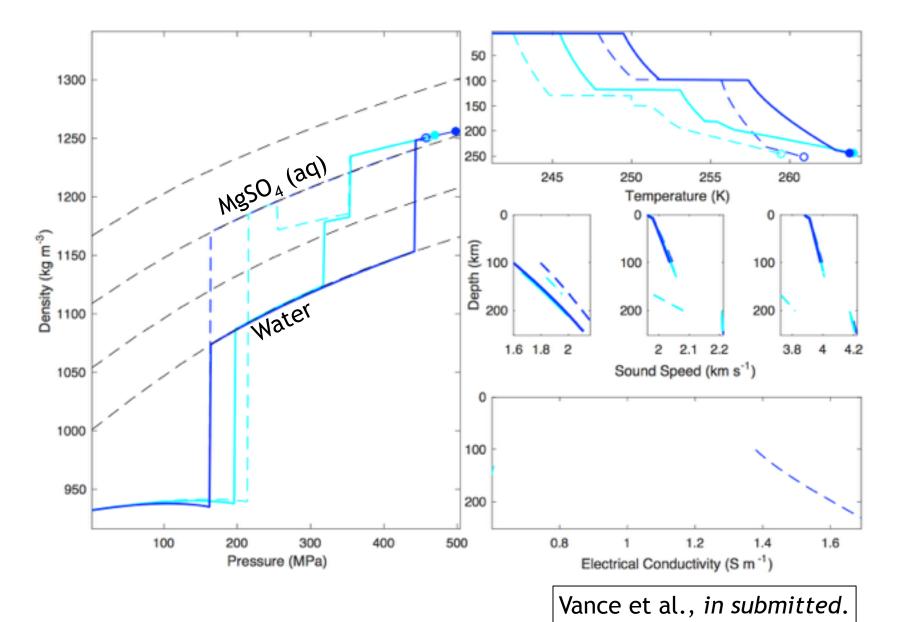


## Enceladus



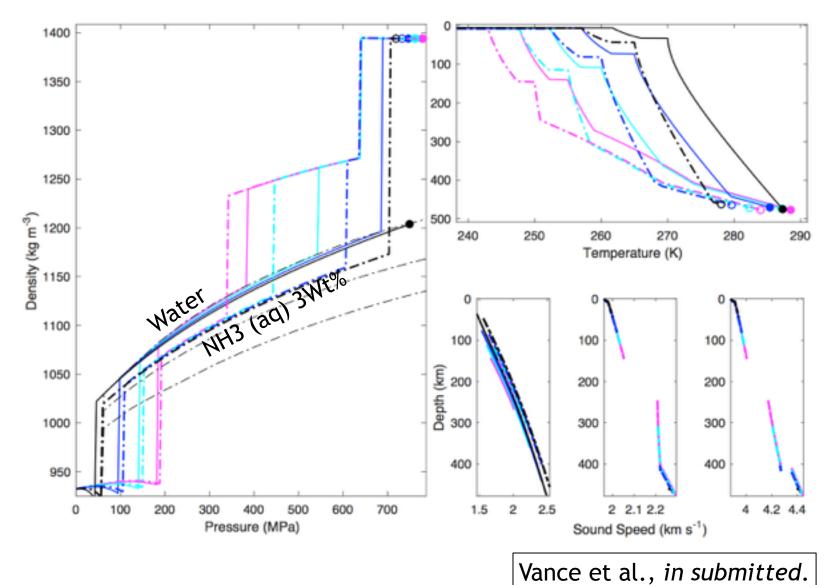
## Callisto



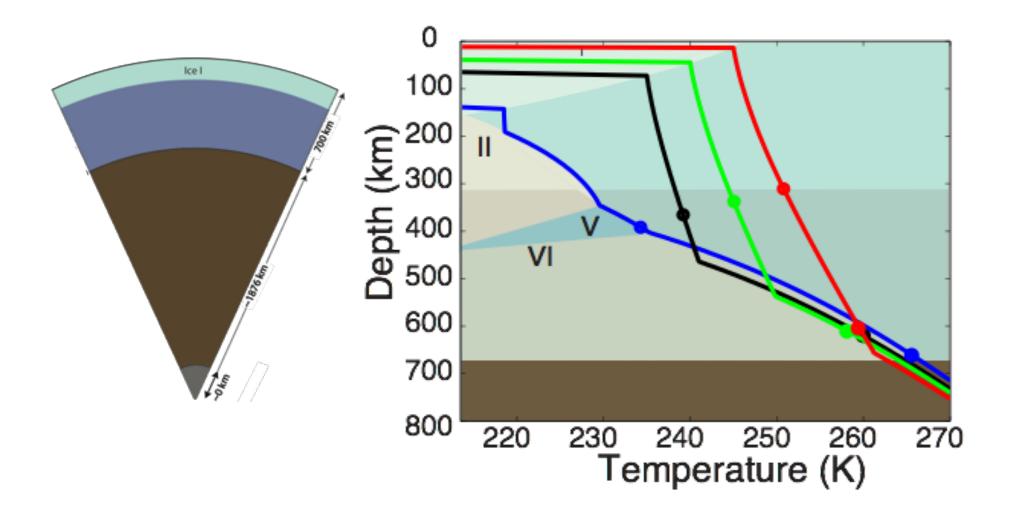


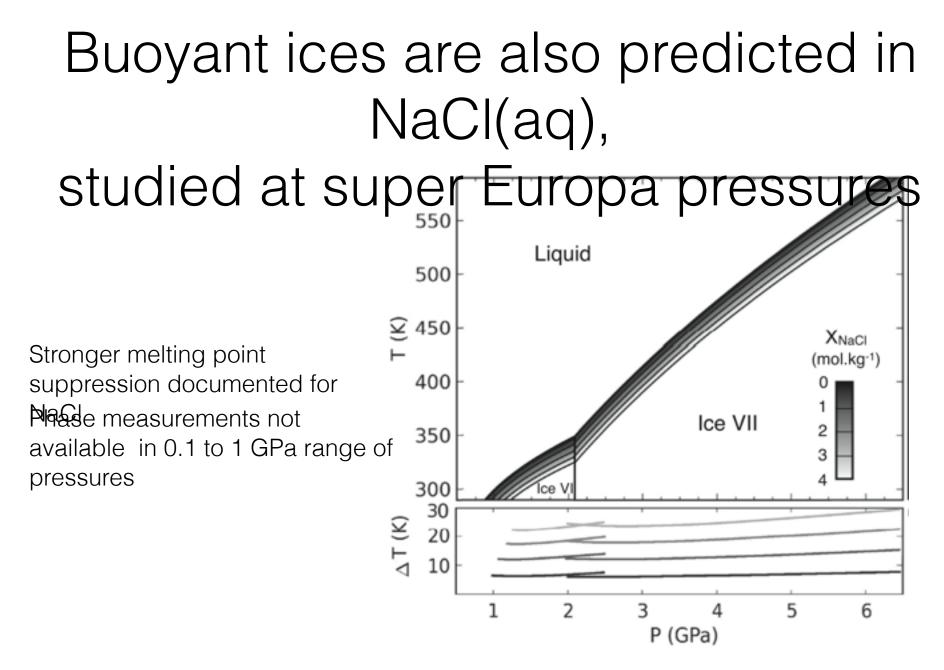
## Titan





## Titan: 15 wt% NH<sub>3</sub> ocean





Journaux+ 2013

#### Buoyant ices are also predicted in NaCl(aq), studied at super Europa pressures High density at low ice VII temperature 1500 18 Wt % 1400 ice VI Density (kg m<sup>-3</sup>) Thermodynamics not available near freezing point 6 Wt % (T<293 K) $(1 \text{ mol } \text{kg}^{-1})$ T = 293 K T = 373 K 1100 T = 473 K T = 573 K 1000 T = 673 K 900 2 з 4 Pressure (GPa) NaCl data from Mantegazzi+2013 Ice densities from Reimers and Watts 1984, Vega+ 2005, Choukroun+

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# Geophysical tests for habitability

- Requiring thermodynamic consistency helps to distinguish among solutions to geophysical datasets
- This could enable measurements of ocean pH
- Titan and Callisto may lack high pressure ices
- Thermodynamic consistency requires further development of EOS for ices fluids, and rocks

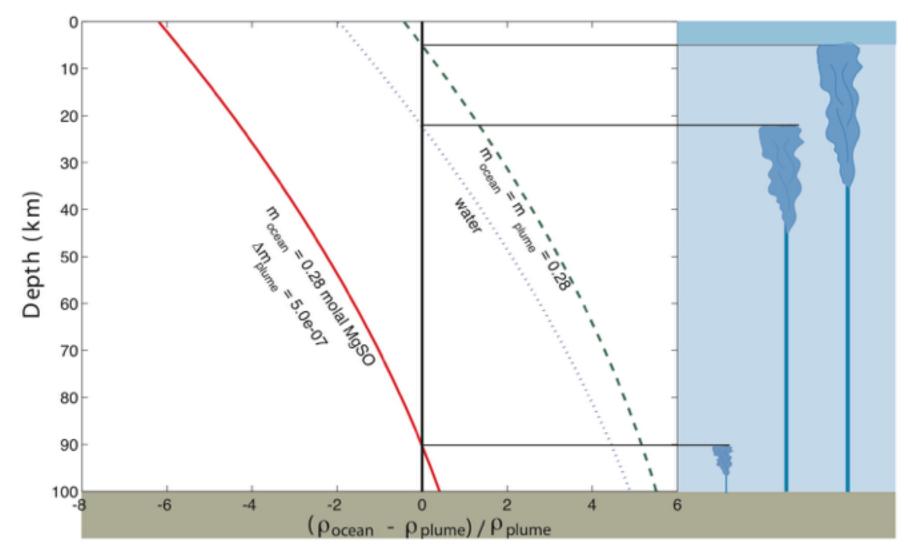




Distance from Sun (temperature)

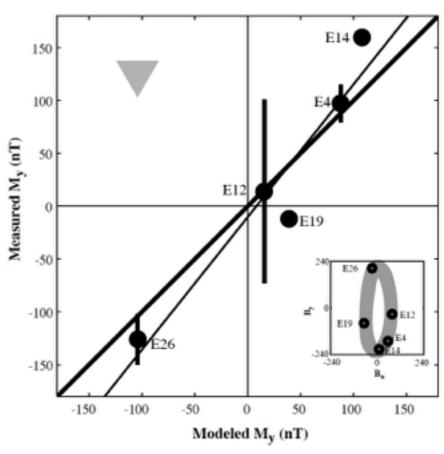


## Oceanic plumes may not always reach the surface



Vance and Brown 2005, Vance and Goodman 2009

# Evidence for present-day oceans in Europa, Ganymede, Callisto



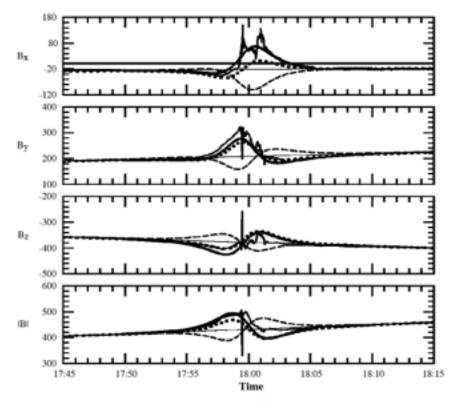


Fig. 4. The E26 pass on 3 January 2000: data within 5 R<sub>g</sub> from Europa's center. Curves are as in Fig. 2. The dashed curves computed from the dipole moment fit to the E4 pass are in antiphase to both the dipole fitted to the E26 data (heavy solid curves) and to the inductive response model (dotted curves).

Kivelson et al. 2000

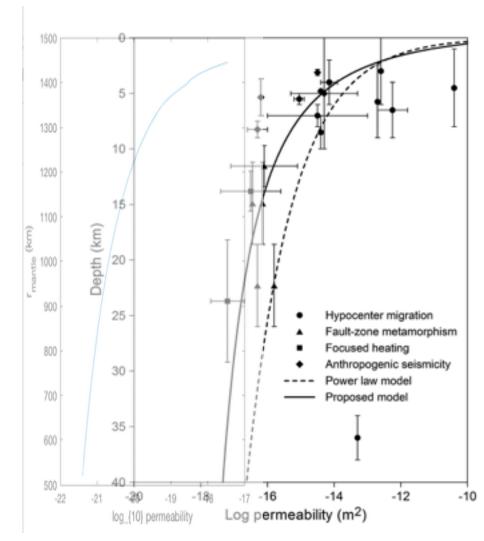
# Porosity and Permeability

$$log(\phi) = -0.65 - 0.1d + 0.0019d^2,$$
  
 $\rho = \phi \rho_f + (1 - \phi) \rho_g,$  Vitovtova et al

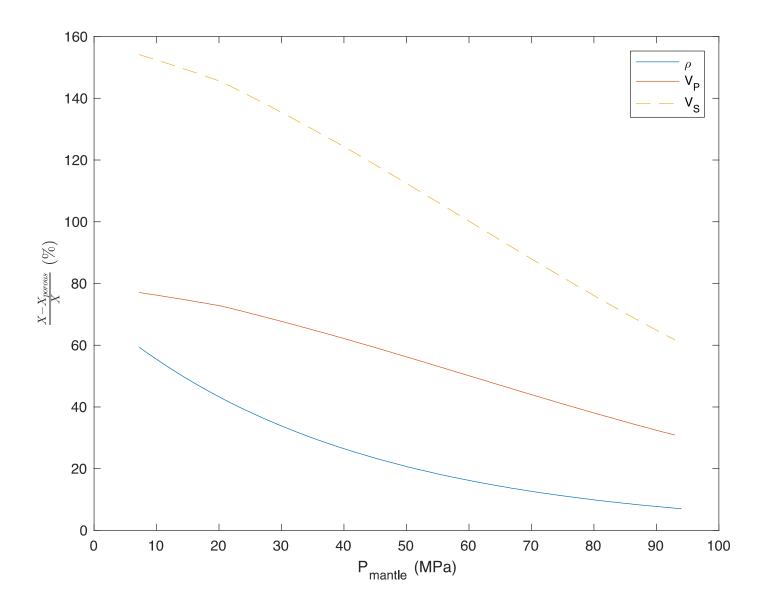
. 2014

 $\log k = \log k_r + (\log k_s - \log k_r) (1+z)^{-\alpha}$ 

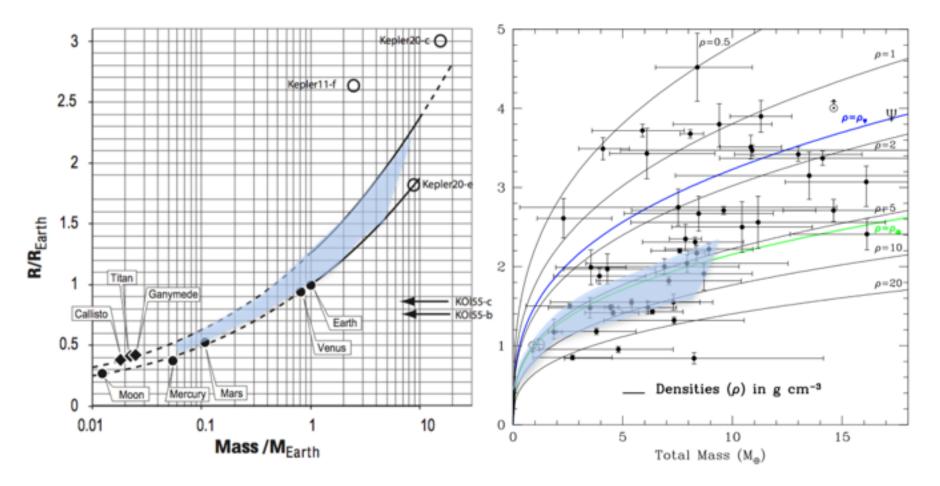
Kuang and Jiao 2014



## Enceladus



### Oceanic exoplanets should define a portion of the M-R curve Additional axes are needed for stellar type, orbit distance...

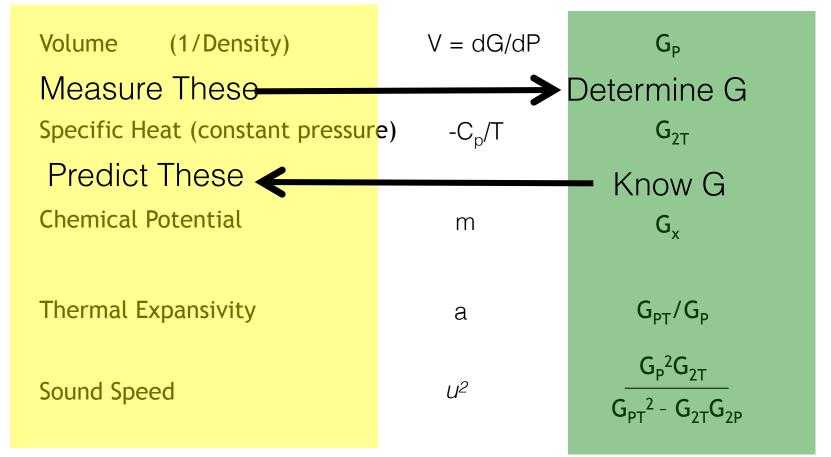


modified from Baraffe+ 2014

modified from Howe+ 2014

## Gibbs Energy Derivatives

G is solution of an ODE



Gibbs energy at high pressure is accurately determined from sound speeds vs P, T, and X

